



Multiscale analysis of tropical cyclones

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Workshop: Multiscale Asymptotics

Asymptotic modelling framework

Structure of atmospheric vortices I: two scales

Structure of atmospheric vortices II: cascade of scales

Conclusions

Nondimensionalization

$$(oldsymbol{x},z) = rac{1}{h_{
m sc}} \left(oldsymbol{x}',z'
ight), \qquad t = rac{u_{
m ref}}{h_{
m sc}} t'$$

$$(\boldsymbol{u}, w) = rac{1}{u_{ ext{ref}}} (\boldsymbol{u}', w'), \qquad (p, T,
ho) = \left(rac{p'}{p_{ ext{ref}}}, rac{T'}{T_{ ext{ref}}}, rac{
ho' RT_{ ext{ref}}}{p_{ ext{ref}}}
ight)$$

where

$$\underline{u_{\rm ref}} = \frac{2}{\pi} \frac{g h_{\rm sc}}{\Omega a} \frac{\Delta \Theta}{T_{\rm ref}} \qquad (\text{thermal wind scaling})$$

Dimensionless numbers, length scales, distinguished limit

Frint $\sim \varepsilon$ $L_{\rm mes} = \varepsilon^{-1} h_{\rm sc}$ ${\rm Ro}_{h_{\rm sc}} \sim \varepsilon^{-1}$ $L_{\rm syn} = \varepsilon^{-2} h_{\rm sc}$ ${\rm Ro}_{L_{\rm Ro}} \sim \varepsilon$ $L_{\rm Ob} = \varepsilon^{-5/2} h_{\rm sc}$ ${\rm Ma} \sim \varepsilon^{3/2}$ $L_{\rm p} = \varepsilon^{-3} h_{\rm sc}$

Compressible flow equations with general source terms

$$\begin{split} &\left(\frac{\partial}{\partial t} + \boldsymbol{v}_{\scriptscriptstyle \parallel} \cdot \nabla_{\scriptscriptstyle \parallel} + w \frac{\partial}{\partial z}\right) \boldsymbol{v}_{\scriptscriptstyle \parallel} + \boldsymbol{\varepsilon} \left(2\boldsymbol{\Omega} \times \boldsymbol{v}\right)_{\scriptscriptstyle \parallel} + \frac{1}{\boldsymbol{\varepsilon}^{3}\rho} \nabla_{\scriptscriptstyle \parallel} p \ = \ \boldsymbol{S}_{\boldsymbol{v}_{\scriptscriptstyle \parallel}}, \\ &\left(\frac{\partial}{\partial t} + \boldsymbol{v}_{\scriptscriptstyle \parallel} \cdot \nabla_{\scriptscriptstyle \parallel} + w \frac{\partial}{\partial z}\right) w \ + \ \boldsymbol{\varepsilon} \left(2\boldsymbol{\Omega} \times \boldsymbol{v}\right)_{\perp} + \frac{1}{\boldsymbol{\varepsilon}^{3}\rho} \frac{\partial p}{\partial z} \ = \ \boldsymbol{S}_{w} - \frac{1}{\boldsymbol{\varepsilon}^{3}}, \\ &\left(\frac{\partial}{\partial t} + \boldsymbol{v}_{\scriptscriptstyle \parallel} \cdot \nabla_{\scriptscriptstyle \parallel} + w \frac{\partial}{\partial z}\right) \rho \ + \ \rho \nabla \cdot \boldsymbol{v} \ = \ 0, \\ &\left(\frac{\partial}{\partial t} + \boldsymbol{v}_{\scriptscriptstyle \parallel} \cdot \nabla_{\scriptscriptstyle \parallel} + w \frac{\partial}{\partial z}\right) \Theta \ = \ \boldsymbol{S}_{\Theta}. \end{split}$$





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Tropical easterly african waves



http://www.aoml.noaa.gov/hrd/tcfaq/A4.html

Vortex tilt in the incipient hurricane stage

(Velocity potential)



Dunkerton et al., Atmos. Chem. Phys., 9, 5587-5646 (2009)

Radial momentum balance regimes

$$-\frac{1}{\rho}\frac{\partial p}{\partial r} + fu_{\theta} = \mathcal{O}(1) \quad \text{geostrophic} \qquad \text{Ro} \ll 1 \qquad \text{typical "weather"}$$
$$\frac{u_{\theta}^{2}}{r} - \frac{1}{\rho}\frac{\partial p}{\partial r} + fu_{\theta} = \mathcal{O}(1) \quad \text{gradient wind} \qquad \text{Ro} = \mathcal{O}(1) \qquad \begin{array}{c} \text{tropical storm} \\ \text{incipient hurricane} \end{array}$$
$$\frac{u_{\theta}^{2}}{r} - \frac{1}{\rho}\frac{\partial p}{\partial r} \qquad = \mathcal{O}(1) \quad \text{cyclostrophic} \qquad \text{Ro} \gg 1 \qquad \begin{array}{c} \text{hurricane} \end{array}$$

Päschke, Marschalik, Owinoh, K., JFM, **701**, 137–170 (2012)

Dörffel et al., preprint, arXiv:1708.07674 (2017)

Asymptotic scaling regime





$$t_{\rm syn} = \frac{h_{\rm sc}/u_{\rm ref}}{\boldsymbol{\varepsilon}^2}; \qquad L_{\rm syn} = \frac{h_{\rm sc}}{\boldsymbol{\varepsilon}^2}; \qquad |\boldsymbol{v}_{\scriptscriptstyle \parallel}| = \mathcal{O}(1);$$

farfield: classical QG theory
 $|\boldsymbol{v}_{\scriptscriptstyle \parallel}| L = \mathcal{O}(\boldsymbol{\varepsilon}^{-2}); \quad |\boldsymbol{v}_{\scriptscriptstyle \parallel}|/fL = \mathcal{O}(\boldsymbol{\varepsilon})$

$$L_{\text{mes}} = \frac{h_{\text{sc}}}{\boldsymbol{\varepsilon}^{3/2}}; \qquad |\boldsymbol{v}_{\text{H}}| = \mathcal{O}\left(\frac{1}{\boldsymbol{\varepsilon}^{1/2}}\right)$$

core: gradient wind scaling
$$|\boldsymbol{v}_{\text{H}}| L = \mathcal{O}\left(\boldsymbol{\varepsilon}^{-2}\right); \quad |\boldsymbol{v}_{\text{H}}|/fL = \mathcal{O}\left(1\right)$$

Result of matched asymptotic expansion analysis:

3D Theory for

vortex motion, vortex core dynamics*,

and the role of subscale moist processes*

* Includes strong vortex tilt

* Modelled by prescribed heating patterns here

Vortex motion

Adiabatic case ($Q_{\Theta} \equiv 0$)



- Linear small displacement* theory extended to large displacements
- Precession and stationary tilt in background shear* explained analytically

Vortex motion

$$\boldsymbol{X}(\tau, z)/L_{\text{mes}} = \overline{\boldsymbol{X}}(\tau)/\sqrt{\boldsymbol{\varepsilon}} + \widehat{\boldsymbol{X}}(\tau, z).$$

where

 $d\overline{X}/d\tau = \overline{v}_{\rm qg}$ $\partial \widehat{\boldsymbol{X}} / \partial \tau = \widehat{\boldsymbol{X}} \cdot (\boldsymbol{\nabla}_{\!\!\!\text{\tiny H}} \overline{\boldsymbol{v}}_{qg}) + \widehat{\boldsymbol{v}}_{qg}^{*} - \underbrace{\left(\ln \frac{1}{\sqrt{\boldsymbol{\varepsilon}}} + \frac{1}{2} \right) (\boldsymbol{k} \times \boldsymbol{\chi})^{*} + (\boldsymbol{k} \times \boldsymbol{\Psi})}_{\boldsymbol{\boldsymbol{\varepsilon}}}$ background advection self-induced motion $\chi = \frac{f^2}{4\pi\overline{\rho}\Gamma} \frac{\partial}{\partial z} \left(\frac{\overline{\rho}\Gamma^2}{d\overline{\Theta}/dz} \frac{\partial \widehat{X}}{\partial z} \right)$ (LIA-like expression) Ψ :

 integral expression representing net vertical transport of horizontal momentum depend on core structure, tilt, diabatic source terms

* includes effect of vortex on background flow (β -gyres)

* Analogous to local-induction-approximation LIA

Adiabatic lifting and WTG

(0th & 1st circumferential Fourier modes: $w = w_0 + w_{11} \cos \theta + w_{12} \sin \theta + ...$)

gradient wind balance (0th) and hydrostatics (1st) in the tilted vortex

$$\frac{1}{\overline{\rho}}\frac{\partial p}{\partial r} = \frac{u_{\theta}^2}{r} + f u_{\theta}, \qquad \Theta_{1\boldsymbol{k}} = -\frac{1}{\overline{\rho}}\frac{\partial p}{\partial r} \frac{\partial}{\partial z} \left(\boldsymbol{e}_r \cdot \widehat{\boldsymbol{X}}\right)_{1\boldsymbol{k}}$$

potential temperature transport (1st)

$$-(-1)^k \frac{u_\theta}{r} \Theta_{1\mathbf{k}^*} + w_{1k} \frac{d\Theta}{dz} = Q_{\Theta,1\mathbf{k}} \qquad (\mathbf{k}^* = 3 - k)$$

1st-mode phase relation: vertical velocity – diabatic sources & vortex tilt

$$\underline{w_{1k}} = \frac{1}{d\overline{\Theta}/dz} \left[\underline{Q_{\Theta,1k}} + \left(\boldsymbol{e}_r \cdot \frac{\partial \widehat{\boldsymbol{X}}^{\perp}}{\underline{\partial z}} \right)_{\boldsymbol{k}} \frac{u_{\theta}}{r} \left(\frac{u_{\theta}^2}{r} + f \, u_{\theta} \right) \right]$$

Spin-up by asynchronous heating

$$\underbrace{\frac{\partial u_{\theta,0}}{\partial \tau} + w_0 \frac{\partial u_{\theta,0}}{\partial z} + u_{r,00} \left(\frac{\partial u_{\theta}}{\partial r} + \frac{u_{\theta}}{r} + f\right)}_{\text{standard axisymmetric balance}} = -\boldsymbol{u_{r,*}} \left(\frac{u_{\theta}}{r} + f\right)$$

$$\boldsymbol{u_{r,*}} = \left\langle w \frac{\partial}{\partial z} \left(\boldsymbol{e_r} \cdot \widehat{\boldsymbol{X}} \right) \right\rangle_{\theta} = \frac{1}{d\overline{\Theta}/dz} \left(Q_{\Theta,11} \frac{\partial \widehat{X}}{\partial z} + Q_{\Theta,12} \frac{\partial \widehat{Y}}{\partial z} \right)$$

$$\boldsymbol{e}_{r} \cdot \widehat{\boldsymbol{X}} = \widehat{X} \cos \theta + \widehat{Y} \sin \theta$$
$$w_{1\boldsymbol{k}} = \frac{1}{d\overline{\Theta}/dz} \left[Q_{\Theta,1\boldsymbol{k}} + \frac{\partial}{\partial z} \left(\boldsymbol{e}_{r} \cdot \widehat{\boldsymbol{X}}^{\perp} \right)_{\boldsymbol{k}} \frac{u_{\theta}}{r} \left(\frac{u_{\theta}^{2}}{r} + f u_{\theta} \right) \right]$$

Spin-up by asynchronous heating

$$\underbrace{\frac{\partial u_{\theta,0}}{\partial \tau} + w_0 \frac{\partial u_{\theta,0}}{\partial z} + u_{r,00} \left(\frac{\partial u_{\theta}}{\partial r} + \frac{u_{\theta}}{r} + f\right)}_{\text{standard axisymmetric balance}} = - \boldsymbol{u_{r,*}} \left(\frac{u_{\theta}}{r} + f\right)$$



Apparent radial transport in a tilted vortex

Theory is valid uniformly for large vortex Rossby numbers $(f \rightarrow 0)$ as long as the internal wave Froude number is small

20.0 17.5 50 · 15.0 s/m ui $\frac{|\theta_n|}{|\theta_n|}$ max max $\frac{|\theta_n|}{10.0}$ 0 Y in km -50 -1005.0 -150 · 2.5 0.0 ہٰ X in km -100 -50 50 100 20 40 60 80 100 120 0 140 time in h Centerline evolution Intensification

Qualitative corroboration through 3D-numerics (benign case)

$$w_{1k} = \frac{1}{d\overline{\Theta}/dz} \left[-\frac{\partial}{\partial z} \left(\boldsymbol{e}_r \cdot \widehat{\boldsymbol{X}} \right)_{\boldsymbol{k}} \frac{u_{\theta}^0}{r} \left(\frac{u_{\theta}^{02}}{r} + f \, u_{\theta}^0 \right) + \frac{\partial}{\partial z} \left(\boldsymbol{e}_r \cdot \widehat{\boldsymbol{X}}^{\perp} \right)_{\boldsymbol{k}} \frac{u_{\theta}}{r} \left(\frac{u_{\theta}^2}{r} + f \, u_{\theta} \right) \right]$$

Qualitative corroboration through 3D-numerics (violent case)*



$$w_{1k} = \frac{1}{d\overline{\Theta}/dz} \left[-\frac{\partial}{\partial z} \left(\boldsymbol{e}_r \cdot \widehat{\boldsymbol{X}} \right)_{\boldsymbol{k}} \frac{u_{\theta}}{r} \left(\frac{u_{\theta}^2}{r} + f \, u_{\theta} \right) + \frac{\partial}{\partial z} \left(\boldsymbol{e}_r \cdot \widehat{\boldsymbol{X}}^{\perp} \right)_{\boldsymbol{k}} \frac{u_{\theta}}{r} \left(\frac{u_{\theta}^2}{r} + f \, u_{\theta} \right) \right]$$

* Ultimately leaves asymptotic regime!

Compatibility with Lorenz' APE theory

$$\left(re_{\mathbf{k}}\right)_{t} + \left(ru_{r,0}[e_{\mathbf{k}}+p']\right)_{r} + \left(rw_{0}[e_{\mathbf{k}}+p']\right)_{z} = \frac{r\overline{\rho}}{N^{2}\overline{\Theta}^{2}}\left(\Theta_{0}'Q_{\Theta,0} + \Theta_{1}'\cdot Q_{\Theta,1}\right)$$

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Cascade of scales:

radial structure

- > 2 radial layers (eye)
- vortex Rossby waves
- spiral rainbands
- "spotty" cloud patterns



Cascade of scales:

vertical structure

- boundary layer
- convective updrafts
- secondary circulation
- tropopause cap



Boundary layer

• turbulent friction disturbs momentum balance

$$-\frac{u_{\theta}^{2}}{r} - f u_{\theta} + \frac{1}{\overline{\rho}}\frac{\partial p}{\partial r} = \frac{\partial}{\partial z}\left(K\frac{\partial u_{r}}{\partial z}\right)$$
$$-2\frac{u_{\theta}u_{r}}{r} + f u_{r} = \frac{\partial}{\partial z}\left(K\frac{\partial u_{\theta}}{\partial z}\right)$$

- implies Ekman-type radial inflow
- boundary layer expells mass vertically

Orders of magnitude:

 $u_{r} \sim u_{\theta} \sim \frac{u_{\text{ref}}}{\delta}; \qquad L \sim \frac{h_{\text{sc}}}{\delta^{2}}; \qquad h_{\text{bl}} \sim \delta^{2} h_{\text{sc}}; \qquad w \sim \delta^{\alpha} u_{\text{ref}}$ $2\pi \left(\frac{h_{\text{sc}}}{\delta^{2}}\right) \left(\frac{u_{\text{ref}}}{\delta}\right) \sim \pi \left(\frac{h_{\text{sc}}}{\delta^{2}}\right)^{2} \delta^{\alpha} u_{\text{ref}} \qquad \Rightarrow \qquad w_{\text{bl-top}} \sim \delta^{3} u_{\text{ref}}$

Same order of magnitude as observed in the tilted bulk vortex



Convective updrafts



Orders of magnitude

$$w_{\rm upd} \le \sqrt{2\,{\rm CAPE}}$$

CAPE ~ 100 ... 3 600 m²/s²;
$$w_{\rm upd} \sim 10 \dots 60 \text{ m/s} \sim \frac{u_{\rm ref}}{\delta^{\beta}}; \beta \geq 0$$

vertical velocities

$$w_{\rm bulk} \sim w_{\rm bl-top} \ll w_{\rm upd}$$

but vertical mass fluxes

 $\dot{m}_{\text{bulk}} \sim \dot{m}_{\text{bl-top}} \sim \dot{m}_{\text{upd}} \qquad \Rightarrow \qquad \boldsymbol{w}_{\text{bulk}} \sim \boldsymbol{w}_{\text{bl-top}} \sim \overline{\boldsymbol{w}}_{\text{upd}}$

Convective updrafts



Convection concentrates in narrow towers (area fraction $\sigma \ll 1$) Dry dynamics between towers Comparable average vertical mass fluxes

From angular momentum conservation for a centered torus ...



From angular momentum conservation for a centered torus ... Spin-up by asynchronous convection

$$\underbrace{\frac{\partial u_{\theta,0}}{\partial \tau} + w_0 \frac{\partial u_{\theta,0}}{\partial z} + u_{r,00} \left(\frac{\partial u_{\theta}}{\partial r} + \frac{u_{\theta}}{r} + f\right)}_{\text{standard axisymmetric balance}} = - \boldsymbol{u_{r,*}} \left(\frac{u_{\theta}}{r} + f\right)$$

$$\boldsymbol{u}_{\boldsymbol{r},*} = \left\langle \boldsymbol{w} \frac{\partial}{\partial z} \left(\boldsymbol{e}_r \cdot \widehat{\boldsymbol{X}} \right) \right\rangle_{\theta} = \underline{\boldsymbol{w}}_{\mathrm{upd},11} \frac{\partial \widehat{X}}{\partial z} + \overline{\boldsymbol{w}}_{\mathrm{upd},12} \frac{\partial \widehat{Y}}{\partial z} \qquad \boldsymbol{!}$$

Area averaged updraft fluxes take role of WTG-induced vertical velocities in the dry vortex theory

Convection/heating arrangement for most rapid intensification



- max efficiency for $w_{\rm upd} > \delta u_{\rm ref}$
- "decorrelation" by circumferential advection for $w_{\rm upd} \leq \delta u_{\rm ref}$

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10 Vortex Dominated Flows

Ting • Klein • Knio

Addressed to both graduate students and researchers this monograph provides in-depth analyses of vortex-dominated flows via matched and multiscale asymptotics, and it demonstrates how insight gained through these analyses can be exploited in the construction of robust, efficient, and accurate numerical techniques. The dynamics of slender vortex filaments is discussed in detail, including fundamental derivations, compressible core structure, weakly nonlinear limit regimes, and associated numerical methods. Similarly, the volume covers asymptotic analysis and computational techniques for weakly compressible flows involving vortex-generated sound and thermoacoustics. Applied Mathematical Sciences 161 Omar M. Knio

Vortex Dominated Flows Analysis and Computation for Multiple Scale Phenomena







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